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# Magnetic properties of early Pliocene sediments from IODP Site U1467 (Maldives platform) reveal changes in the monsoon system

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ABSTRACT

We report a study of the magnetic stratigraphy and the anisotropy of isothermal remanent magnetization of Pliocene sediments from International Ocean Discovery Program (IODP) Site U1467 drilled in the Maldives

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Monsoon

platform (Indian Ocean) during Exp. 359. Magnetic stratigraphy gives a precise record of geomagnetic reversals of the early Pliocene from approximately 5.3 Ma to 3.1 Ma providing a detailed age model in an interval where the biostratigraphic record is scarce. We use the anisotropy of isothermal remanent magnetization (AIRM) to investigate the statistical orientation of fine magnetic particles and provide data on the strength and direction of bottom currents during the early Pliocene. The strength of bottom currents recorded by the AIRM, shows a prominent increase at the top of Chron C3n.1n (about 4.2 Ma), and the current direction (NE - SW) is consistent with that of modern instrumental measurements. Since bottom currents in the Maldives are driven by the monsoon, we speculate that the 4.2 Ma increase of bottom currents could mark the onset of the present-day setting, probably related to the coeval uplift phase of the Himalayan plateau.

## 1. Introduction

Wind-induced currents are an important factor controlling the sedimentation in the Maldives archipelago where they regulate sediment transport from the atoll to the deeper part of the platform as well as the geometry of the sedimentary bodies (e.g., Betzler et al., 2009; Lüdmann et al., 2013; Betzler et al., 2016b, Lüdmann et al., 2018). The onset of these current-driven drift deposits has been related to the increase of monsoon activity starting in the middle Miocene (Betzler et al., 2016b). In this paper we report results from paleo and rock-magnetic analyses from International Ocean Discovery Program (IODP) Site U1467 aiming to improve the Pliocene age model and investigate the variations of bottom current strength.

Site U1467 (4° 51.0155′ N and 73° 17.0204′ E) was drilled in the Inner Sea of the Maldives archipelago (Indian Ocean) during IODP Expedition 359 at a water depth of 487.4 m (Betzler et al., 2017) in a distal position compared to the platform margins and moats (Fig. 1). Site U1467 showed nearly horizontally layered seismic reflections of the so-called drift sequences (Betzler et al., 2013; Lüdmann et al., 2013; Wunsch et al., 2017) with no truncations and no indications of mass wasting from the adjacent platform margin. Shipboard analysis suggested that Site U1467 provided a complete and undisturbed succession of all drift sequences from the Late Miocene (Betzler et al., 2017) with good potential for paleoceanographic and paleoclimatic studies. In particular the drift succession, from mid-Miocene to recent, contains several sequences that are potentially related to fluctuations in the monsoon-driven current system. Dating these sequences can yield the ages of changes in the strength and direction of the currents.

The paleomagnetic analysis in this study provides the magnetostratigraphic age of the upper portion of Site U1467 that was sampled with advanced piston coring (APC), and uses the anisotropy of the isothermal remanent magnetization (AIRM) to investigate the statistical orientation of fine magnetic particles aiming to provide a sedimentological record of direction and strength of bottom currents.

Shipboard paleomagnetic measurements of Site U1467 (Betzler et al., 2016a, Betzler et al., 2017) gave poor results because of a combination of two factors: (i) the very low concentration of magnetic minerals in carbonate platform sediments, which resulted in a very weak natural remanent magnetization (NRM) and (ii) a strong magnetic contamination of the cores due to metallic particles presumably originating from the drilling pipes. The contamination covered the original weak paleomagnetic signal of the sediment preventing any valuable

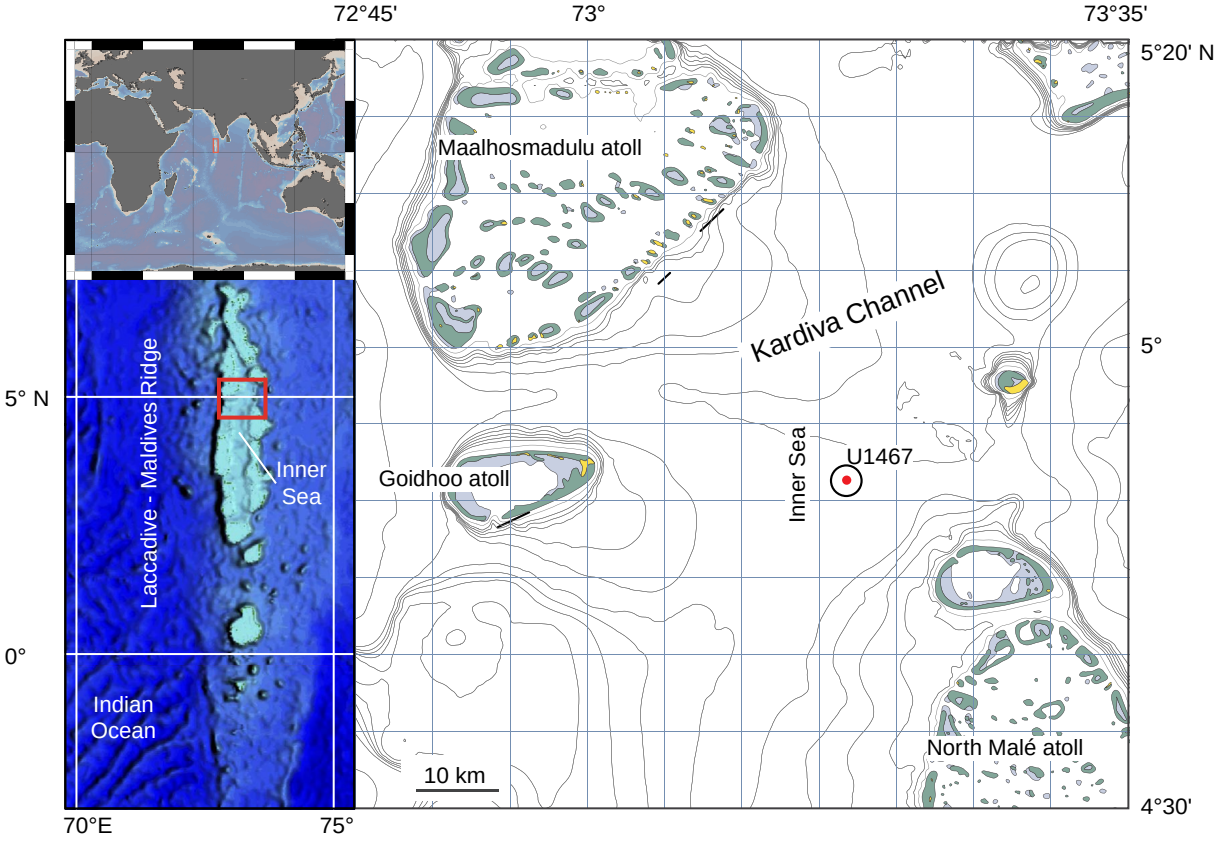



**Fig. 1.** Location map of IODP Site U1467.

measurements with the shipboard pass-through technique that measures half-cores. Measurements of individual 7 $cm^3$ box-samples, taken from the inner part of the core, performed aboard, did not show signs of significant contamination suggesting that it was restricted to the outer part of the cores. Unfortunately, the NRM of these specimens, ranging from ca. $1 \times 10^{-5}$ A/m to $1 \times 10^{-4}$ A/m, was too weak to be measured reliably using the JR-6 spinner magnetometer available on the Joides Resolution. However, the NRM of the supposedly uncontaminated box-samples is within the range of sensitivity of a DC-SQUIDS cryogenic magnetometer, and this provided the motivation to collect and measure 580 standard sediment specimens with the aim to obtain a reliable magnetic stratigraphy of Site U1467.

## 2. Material and sampling

Standard paleomagnetic specimens (Natsuhara-Giken sampling cubes, with a volume of 7 $cm^3$) were collected in the upper part of Site U1467 from core sections 359-U1467B-11H to 359-U1467B-34H and from 359-U1467C-10H to 359-U1467C-17H, corresponding to 84 m to 302 m core depth below sea floor (CSF-A), at the Gulf Coast Repository at Texas A&M University. Specimens were sampled only from azimuthally-oriented APC cores, since information on core orientation is essential for paleomagnetic studies at equatorial latitudes such as for Site U1467.

According to the shipboard sedimentology, the studied part of Site U1467 was divided into three main lithostratigraphic units (Betzler et al., 2017). The uppermost Unit I was recovered in the top 110 m CSF-A and consists of unlithified, foraminifer-rich wackestone to packstone with a predominance of very fine- to fine-grained wackestone. Unit II extends from ca. 110 m to 215 m CSF-A; it comprises late Pliocene sediments and is characterized by interlayered unlithified and partially lithified planktonic foraminifera–rich wackestone and mudstone with pteropods and particulate organic matter. Unit III, which extends from 215 m to 303 m CSF-A, consists of partially lithified very fine–grained mudstone to wackestone with a dominance of wackestone. The sediment contains abundant planktonic foraminifera; echinoid spines and sponge spicules are common while benthic foraminifera are rare. Microfossil preservation throughout Units III and most of Unit II was generally poor to moderate, and in particular the interval from ca. 150 m to 300 m yields no biostratigraphic ages. The interval chosen for paleomagnetic study was also intended to cover this interval with poor or absent biostratigraphy.

## 3. Paleomagnetic analysis

### 3.1. Isothermal remanent magnetization

Magnetic mineralogy was investigated by acquisition of isothermal remanent magnetization (IRM) in a set of pilot specimens. IRM was acquired in 12 stepwise increasing fields from 0.03 T to 1 T, induced using a ASC pulse magnetizer (Fig. 2a). Results indicate that all measured specimens are characterized by the presence of only low-coercivity magnetic minerals that saturate in fields between 100 mT and 250 mT. These coercivities are below the maximum coercivity of uniaxial magnetite (e.g. Tauxe, 2002) and suggest a rather homogeneous mineralogy made of ferromagnetic minerals such as magnetite (or maghemite) without any significant presence of diagenetic iron sulphides that can be distinguished from their higher coercivity (e.g., Tauxe, 2002).

The saturated isothermal remanent magnetization is relatively weak, as often found in carbonate sediments, because of the low concentration of ferrimagnetic minerals. It ranges from $1.0 \times 10^{-3}$ to $3.7 \times 10^{-2}$ A/m, suggesting a wide variability in the concentration of magnetic minerals.

### 3.2. Natural remanent magnetization

The pass-through shipboard measurements of the NRM showed a huge scope of values with intensity ranging from $5 \times 10^{-6}$ A/m to $1 \times 10^{-1}$ A/m along the same core and with a strong downcore decreasing trend (Betzler et al., 2017). This was interpreted as the consequence of steel contamination most likely originating from worn off drill pipes. The bottom part of each APC core, which were the least contaminated, exhibited NRM values ranging from ca. $5 \times 10^{-6}$ A/m to $1 \times 10^{-4}$ A/m that were considered reasonable values for carbonate sediments and thus regarded as uncontaminated or only slightly

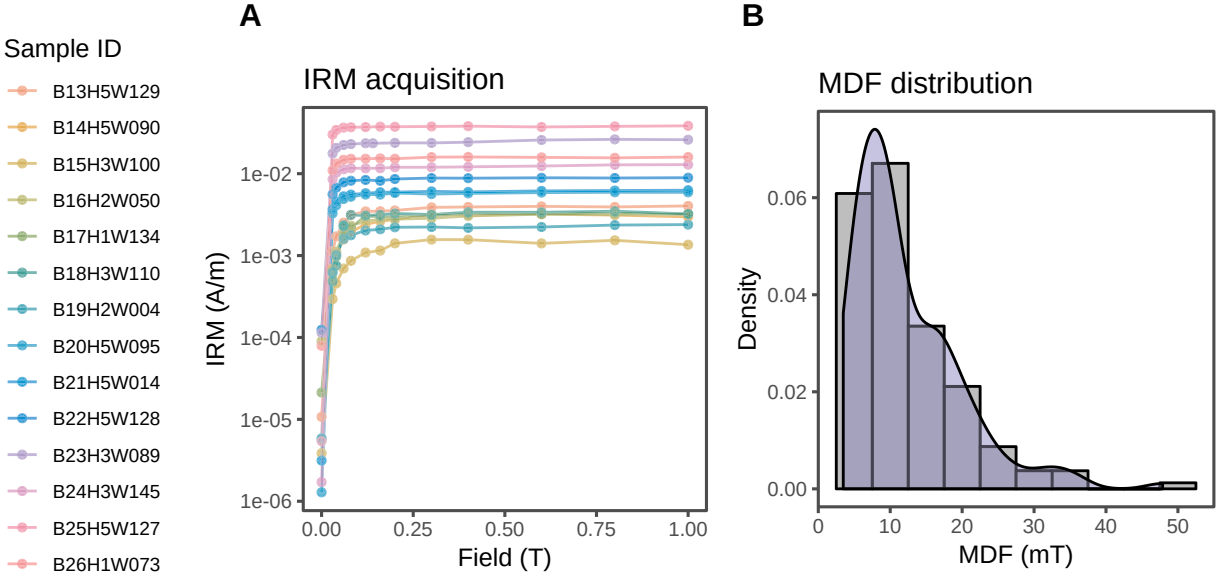



**Fig. 2.** Acquisition of isothermal remanent magnetization of representative samples from the investigated site (A) and the estimate of the density distribution of median destructive field of the natural remanent magnetization (B). Isothermal remanent magnetization acquisition shows that all samples are saturated at fields higher than 100–150 mT indicating the presence of low coercivity minerals. The low coercivity rules out the presence of relevant amounts of hematite or diagenetic iron-sulphides and suggests that magnetite (or maghemite) is the main magnetic mineral in the sediments. The histogram and the density distribution of the median destructive field has a mode of about 10 mT confirming that the natural remanent magnetization is carried by low-coercivity minerals.

contaminated. Discrete specimens taken from the inner part of the cores showed similar NRM intensity values corroborating the hypothesis that contamination was limited to the outer part of the cores.

Based on these remarks, box specimens for the shore-based analysis were collected from the inner part of the core in Site U1467 and were measured using a 2G-enterprise DC-SQUID magnetometers at the CIMaN-ALP laboratory (Cuneo, Italy). Samples were progressively demagnetized in alternating field (AF) up to the maximum field of 100 mT according to a standard paleomagnetic procedure.

The directional components of the natural magnetization were calculated using the method of the principal component analysis (Kirschvink, 1980) and the PuffinPlot software (Lurcock and Wilson, 2012). The quality of the measurements and the line fitting was checked by visual inspection of the orthogonal vector plots and was quantified using the maximum angular deviation (MAD).

The AF demagnetization technique was effective in demagnetizing the NRM testifying that low-coercivity minerals are the main carriers of the NRM. Vector plots generally show a small viscous overprint removed in fields smaller than 20 mT, while the remaining part of the NRM is demagnetized in the field interval from 20 mT to 100 mT. The component isolated within this coercivity interval was used to calculate the characteristic remanent magnetization (ChRM) when sufficiently linear and well defined. However, the success rate in recovering reliable paleomagnetic directions was generally low and many specimens were discarded because they did not yield results with acceptable quality. In specimens that gave acceptable results, on average, about 20% of the NRM was removed after AF demagnetization at 20 mT, and ca. 95% of the NRM was removed at field of 60 mT. Moreover, the NRM median destructive field of acceptable specimens (Fig. 2b) has a modal value of 10 mT. The percentage of NRM removed at 60 mT and the values of median destructive field corroborate the results of IRM acquisition suggesting that stable NRM is carried by pseudo-single domain magnetite. In acceptable specimens, the NRM intensity has an average value of ca. $2.1 \times 10^{-4}$ A/m, the average value of MAD obtained from the vector analysis of these specimens was 8.9°. About 10% of specimens have MAD values between 15° and 20°, which although large, are considered acceptable; most of these specimens are located in the upper part of the investigated interval, between 100 and 180 m depth, or at polarity transitions. Representative orthogonal vector plots for Site U1467 are illustrated in Fig. 3.

The azimuthal orientation of cores was essential for interpretation of magnetic polarity because the paleomagnetic inclinations of equatorial localities, such as Site U1467 during the Miocene, are very close to zero for both normal and reversed polarities; hence the geomagnetic polarities are indistinguishable if based only on inclination data. APC cores collected from Site U1467 were oriented using the "tensor tool" that provided a good first-order orientation. The average declinations from paleomagnetic measurements showed significant departures from North and discrepancies between cores, suggesting that orientation errors of the tensor tool can be as large as ± 30°. However, although large, these errors did not compromise the polarity of ChRM and there was no ambiguity in establishing the magnetic polarity. We did not attempt to remove orientation errors by adjusting the magnetic declination to a mean direction even if this resulted in a reduced precision of the latitude of the virtual geomagnetic pole (VGP).

### 3.3. Anisotropy of isothermal remanent magnetization

In agreement with shipboard measurements, the magnetic susceptibility of box-samples shows negative (diamagnetic) susceptibility, evidence of the dominating diamagnetic matrix of $CaCO_3$ on the ferrimagnetic component. The very weak diamagnetic susceptibility of the

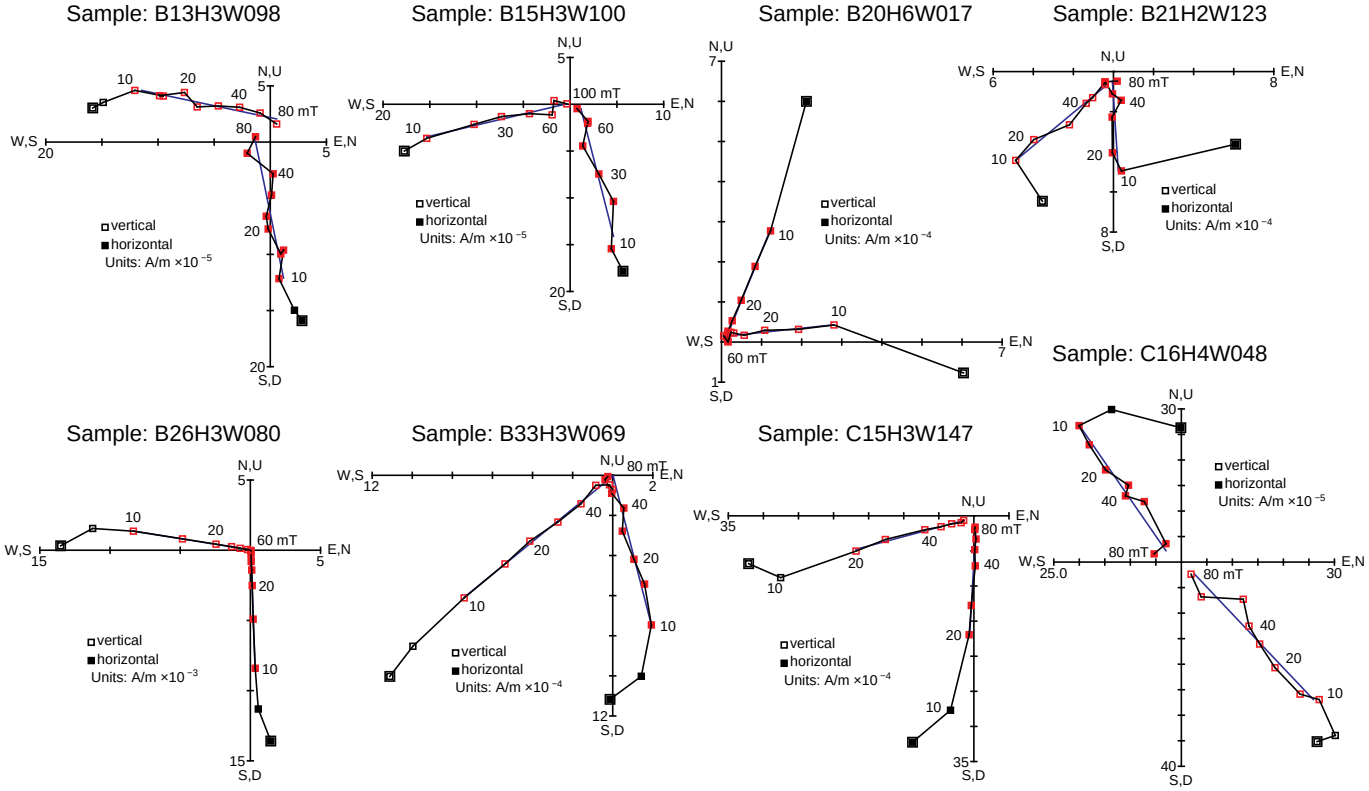



**Fig. 3.** Representative examples of vector plots of alternating field demagnetization of natural remanent magnetization. Stepwise demagnetization of natural remanent magnetization of sediments from Site U1467 shows generally a very small overprint, which is removed at a maximum field of 10–20 mT, followed by a linear path toward the origin that is interpreted as the characteristic remanent magnetization. Blue segments represent the direction of the characteristic remanent magnetization computed as the best-fit line of the selected demagnetization steps (shown in red). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

carbonate sediments recovered in Site U1467 (Betzler et al., 2017) limits the possibility of using the anisotropy of magnetic susceptibility to investigate the orientation pattern of the magnetic particles. Therefore, we resorted to using AIRM which can be measured precisely even in these weakly magnetic sediments.

AIRM measurements were performed in a subset of 75 specimens taken from Core 13 to Core 26 of Hole U1467A. To compute the AIRM an isothermal remanent magnetization induced with a field of 20 mT was measured and then AF demagnetized, repeating this procedure along 6 different axes. Each axis was measured twice along opposite directions for a total of 12 AIRM measurements in each specimen (e.g., Stephenson et al., 1986; Jackson, 1991; Potter, 2004). The intensity of isothermal magnetization was measured with a JR-6 spinner magnetometer and the specimens were demagnetized after each measurement using a tumbling 2G AF-demagnetizer at a maximum field of 80 mT, before inducing the magnetization in the next direction. The anisotropy tensor and the directions of the principal IRM axis $I_i$ (i.e., the eigenvectors of the AIRM tensor) were computed from the remanent magnetization using the AGICO software Anisoft42. The AIRM is therefore represented as a triaxial ellipsoid, whose principal axes correspond to the directions of maximum, intermediate and minimum IRM ($I_1 < I_2 < I_3$). The anisotropy ellipsoid was described using the corrected anisotropy degree (P′) and the shape parameter (T) computed according to Jelinek (1981)

$$T = (2n_1 - n_2 - n_3)/(n_1 - n_3)$$

$$P' = \sqrt{\exp 2\,(a_1^2 + a_2^2 + a_3^2)}$$

where $n_i = \ln I_i$, $a_i = \ln\left(\frac{I_i}{I'_m}\right)$ and $I'_m = \sqrt[3]{I_1 I_2 I_3}$ with $i = 1, 2, 3$ and are shown in Fig. 4.

The direction of the largest axis of the anisotropy tensor $I_1$ represents the magnetic lineation (the preferred orientation of elongated magnetic particles), and the foliation plane is the plane that contains the $I_1$ and $I_2$ directions, hence is orthogonal to the direction of the smallest axis of the anisotropy tensor $I_3$.

## 4. Results and discussion

### 4.1. Magnetostratigraphy

Declination, inclinations of ChRM and the resulting VGP latitude of Site U1467 are shown in Fig. 5 plotted versus depth m CSF-A, together with the available biostratigraphic events from shipboard analysis (Betzler et al., 2017). The comparison of measured levels and acceptable samples in Fig. 5 indicates the low success rate in finding reliable directions of the ChRM. Sometimes, for instance in specimens taken from Hole U1467C, we could not obtain any acceptable results. In general in the upper part of the Site above 110 m, which comprises the sedimentological Unit I, the paleomagnetic data yielded poor results probably related to the coarser granulometry of the unlithified wackestone sediments. At depths between 100 m and 290 m, the quality of the data was sufficient to obtain a reliable record of polarity reversal, although with variable quality.

In the interval with good data quality (i.e., the central part of the record with smaller MAD), the ChRM inclinations are practically indistinguishable from zero, regardless of the polarity, except for transitional directions corresponding to the time elapsed during the reversal of the geomagnetic field. Averaged paleomagnetic directions (Table 1) indicate an equatorial paleolatitude $\lambda = 0.8°$ ($\lambda_{95}^{+} = 4.4°$; $\lambda_{95}^{-} = -2.8°$) of Site U1467 during the early Pliocene within the precision of the paleomagnetic data, which is in agreement with the paleogeographic reconstructions of Besse and Courtillot (2002) and Torsvik et al. (2012).

The record of polarity reversals identifies 6 normal and 6 reversed magnetic polarity zones that, based on the biostratigraphic framework, have been interpreted as Chrons C2An.2n to C3n.4r (Fig. 6) (Gradstein et al., 2012). This interpretation is manly based on the 4 biostratigraphic events available in the studied section and located in the upper part of the record; however, there is some uncertainty even in the biostratigraphic data since the dates based on Last Occurrence of foraminifera (*Dentoglobigerina altispira* and *Globorotalia margaritae*) show a relatively large discrepancy with that of calcareous nannofossil events (LO *Sphenolithus abies* and LO *Reticulofenestra pseudoumbilicus*). Even

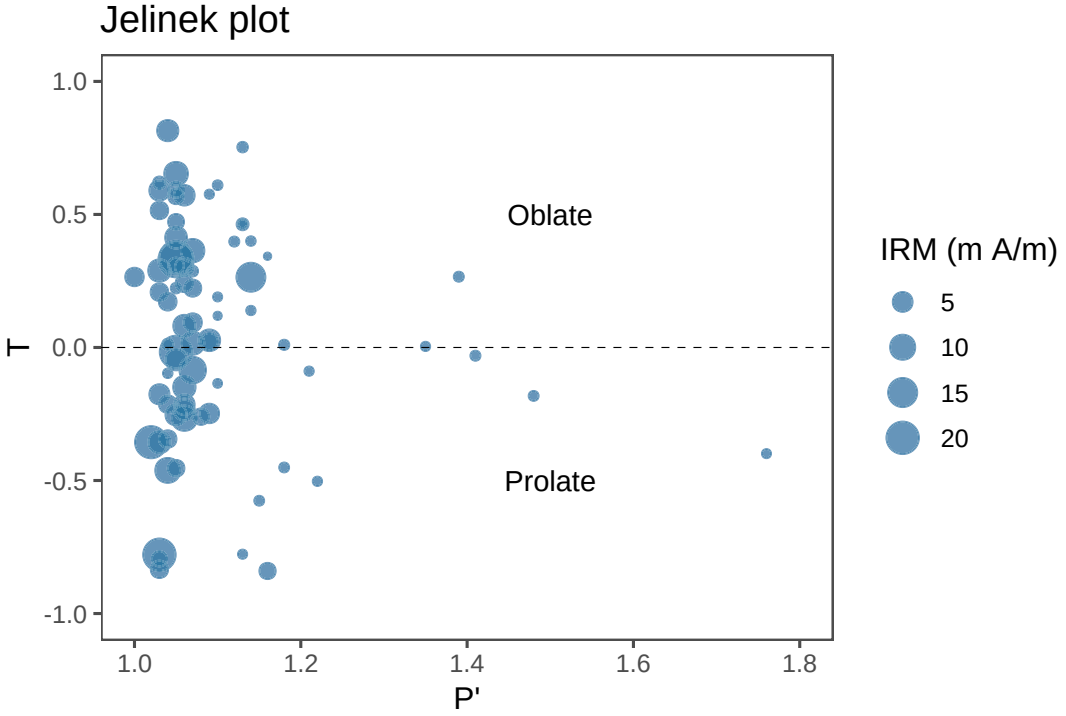



**Fig. 4.** Jelinek plot (Jelinek, 1981) illustrating the shape of anisotropy tensor (T) and corrected degree of anisotropy (P′). Symbol size is proportional to the intensity of isothermal remanent magnetization.

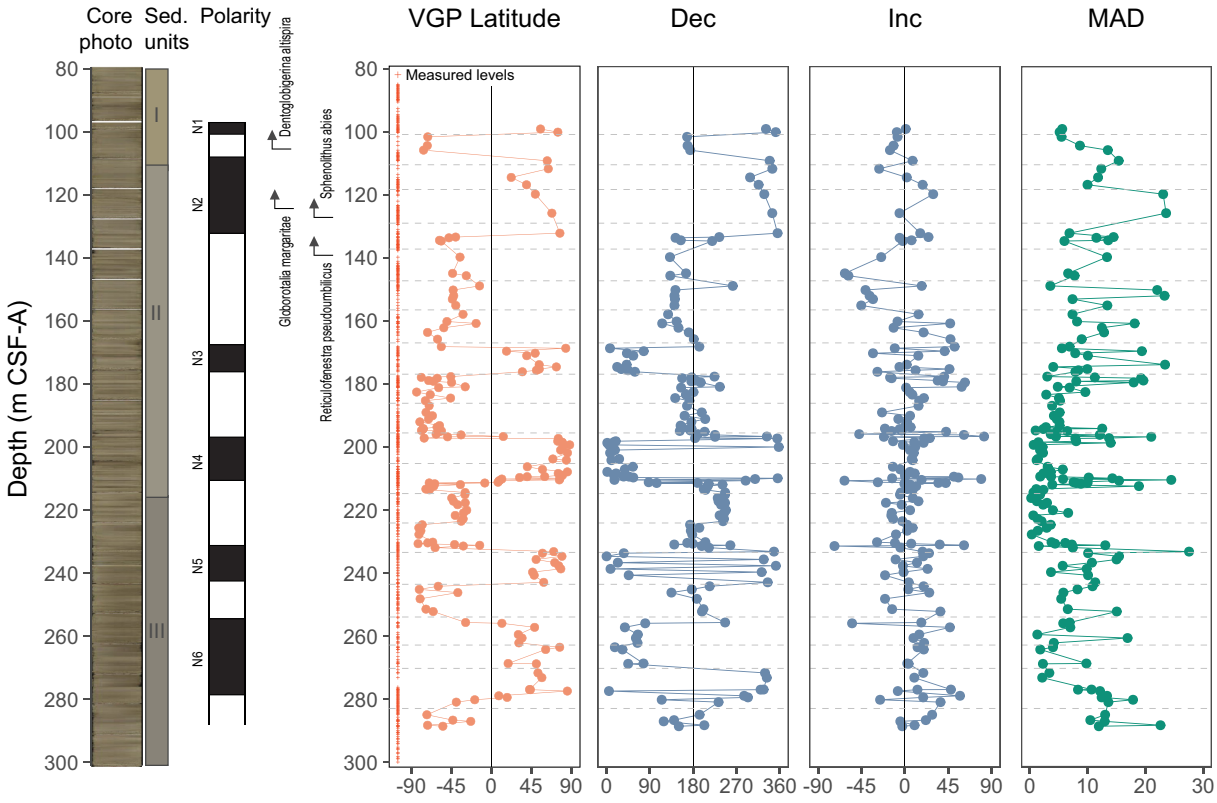



**Fig. 5.** ChRM directions (Declination and Inclination), maximum angular deviation and virtual geomagnetic pole latitude plotted against core depth (m CFS-A). The latitude of the virtual geomagnetic pole is computed from the declination and inclination in order to better interpret the geomagnetic polarities, which are reported in the left column as black and white intervals for normal and reversed polarity, respectively. The horizontal dashed lines indicate cores breaks and the small symbols in the left side of the VGP Latitude panel indicates the measured levels. The biostratigraphic events, core photographs and sedimentary units from Betzler et al. (2017) are also reported. Notice that the paleomagnetic inclinations are not significantly different from zero except for transitional directions, indicating an equatorial paleo-latitude of the site.

**Table 1**
Mean direction, Virtual Geomagnetic Pole position and paleo latitude of Site U1467 (the relatively low precision of the data is mostly a consequence of the poor azimuthal orientation of the cores).

| | |
|---|---|
| Fisher statistics | Dec = 5.7, Inc. = 1.6, $R$ = 121.91, k = 3.12, $a_{95}$ = 7.2, $N$ = 179 |
| VGP | Lat = 82.9, Long = 198.5, $dm_{95}$ = 7.2 $dp_{95}$ = 3.6 |
| Paleo latitude | l = 0.8°, $l_{+95}$ = 4.4°, $l_{-95}$ = −2.8° |

though the few available biostratigraphic markers leave some room in interpreting which magnetochrons correspond to the measured polarity reversals, we believe that our interpretation (Fig. 6) is the best compromise between a reduced variability of the sedimentation rate and the available biostratigraphic framework. Within these limitations, the magnetic stratigraphy of Site U1467 provides a robust age constraint in a section where biostratigraphic records are unavailable. In our interpretation the age of the studied section spans ca. from 5.3 Ma to 3.1 Ma, as reported in detail in Table 2.

### 4.2. Anisotropy of the IRM

According to a number of studies (e.g., Betzler et al., 2009, 2016b, 2018 and references therein) bottom currents in the Maldives platform are considered wind-driven and assumed to be a direct consequence of Asian monsoon. At equatorial latitudes, the link between surface wind and bottom currents extends to a depth of several hundred meters either through Ekman transport or as an undercurrent system and can be seen with modern observations. The present day equatorial Indian Ocean is characterized by seasonally reversing surface currents, known as Wyrtki Jets, driven by zonal winds. Beneath the surface, to a depth of several hundred meters, the flow of the equatorial undercurrent and the equatorial intermediate current has been observed (e.g., Knox, 1976; Reppin et al., 1999; Schott and McCreary, 2001; Iskandar et al., 2009; Nyadjro and McPhaden, 2014). In contrast to other oceans, the Indian Ocean equatorial undercurrent is transient and strongly dependent on winds and pressure gradient variations. Both eastward and westward flows of sub-surface currents have been observed, although modeling studies based on the present day suggest that eastward undercurrents are more likely to occur than westward ones (Schott and McCreary, 2001). Since sub-surface currents develop as consequences of surface wind it is reasonable to assume that stronger surface winds will increase the strength of the undercurrent, and following this argument, we interpret bottom current strength as proxy of the paleo-monsoon.

Magnetic methods are particularly useful when other quick methods for determining paleo flow from sediment beds such as macroscopic paleocurrent indicators (e.g., cross-stratification and sole marks) are lacking. In standard analysis of magnetic grain shape fabric, AIRM is considered to be a proxy for the preferred alignment of elongated natural magnetic particles attained in the final stages of transport, with $I_1$ and $I_3$ representing preferred orientations of the longest and shortest grain axes, respectively (e.g., Hamilton and Rees, 1970, Taira and Peter Scholle, 1979; Novak et al., 2014, Felletti et al., 2016). The method assumes implicitly that the uniaxial shape-anisotropy of magnetic particles dominates triaxial magnetocrystalline anisotropy, as expected for elongated magnetite particles (e.g., Tauxe, 2002).

According to theoretical, experimental and field-based fabric studies, two main anisotropic fabric patterns are found (e.g., Harms et al., 1982; Baas et al., 2007): (i) flow-aligned fabric; and (ii) flow-transverse

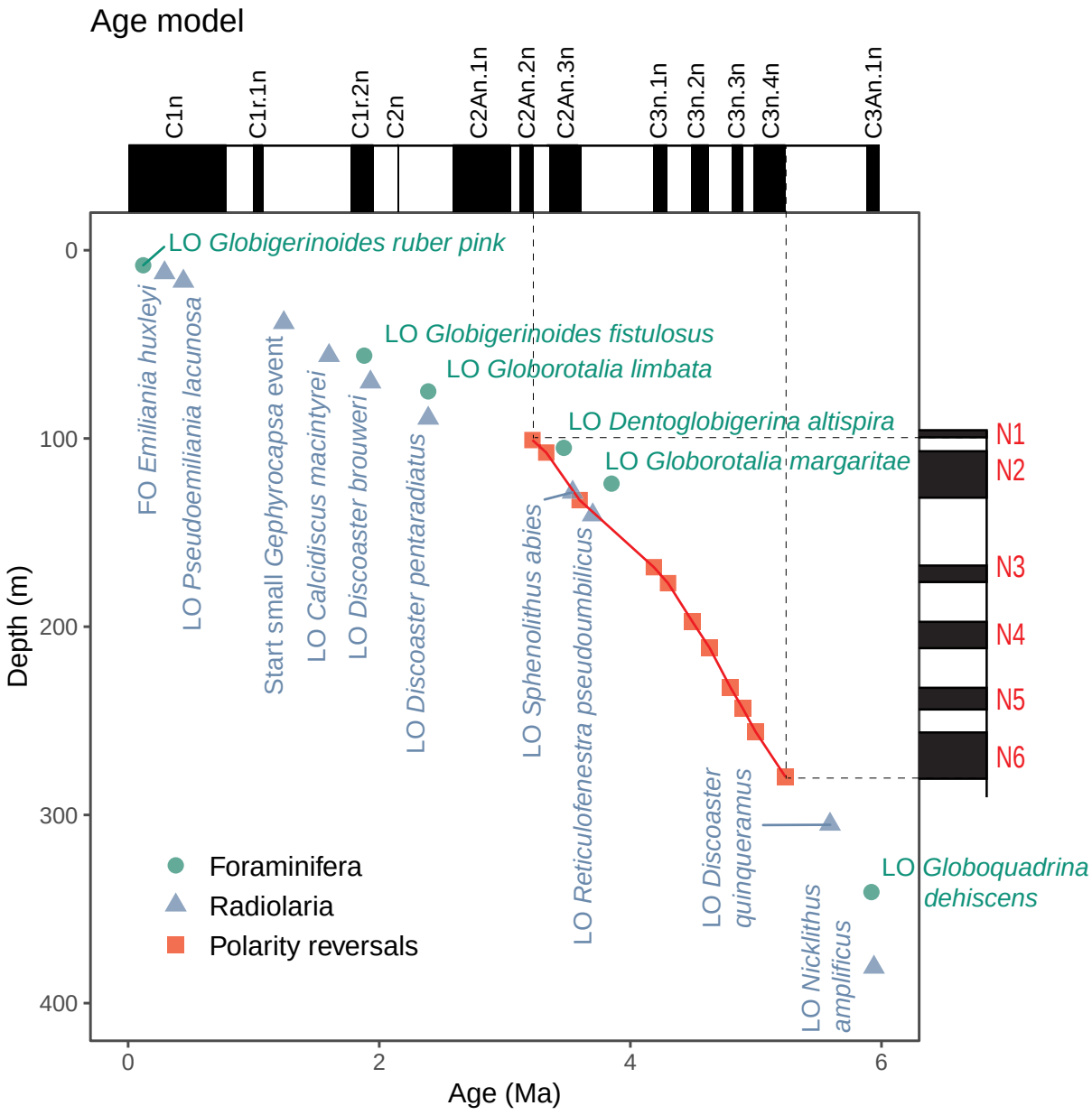
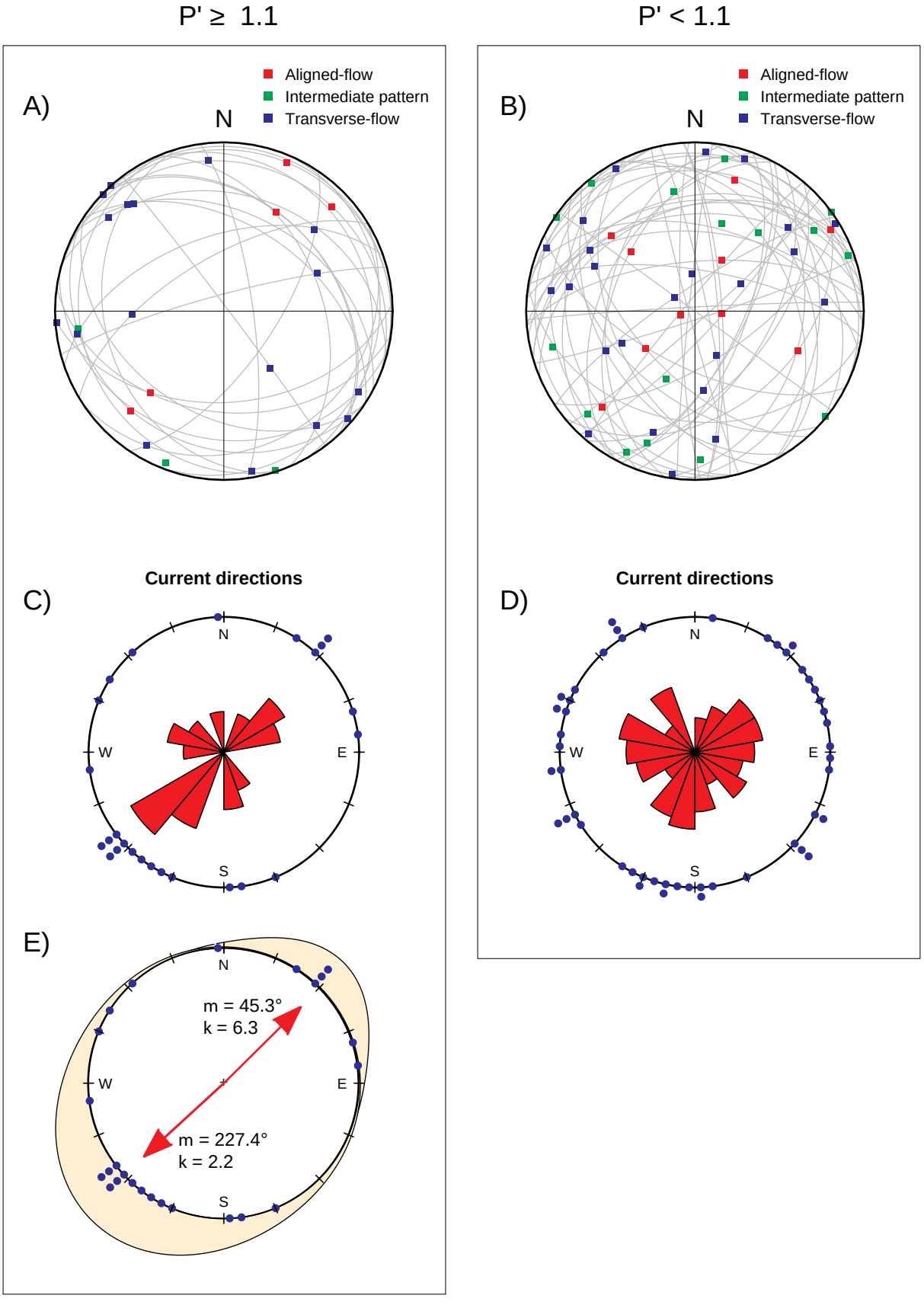



**Fig. 6.** Paleomagnetic interpretation and age model of the studied portion of Site U1467. Shipboard biostratigraphic events are reported to provide the general age frame. The reversal polarity sequence, N1 to N6, from Site U1467 is shown in the right-vertical axis. The open circles connected by the red line represent the correlation of this polarity reversal sequence to the reference geomagnetic polarity scale on the horizontal upper axis. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

**Table 2**
Magnetostratigraphic reversals.

| Chron | Age | Depth top | Depth bottom |
|---|---|---|---|
| | (Ma) | (m CSF-A) | (m CSF-A) |
| C2An.2n bottom | 3.220 | 100.10 | 101.59 |
| C2An.3n top | 3.330 | 105.81 | 109.18 |
| C2An.3n bottom | 3.596 | 132.15 | 133.38 |
| C3n.1.n top | 4.187 | 168.14 | 168.65 |
| C3n.1n bottom | 4.300 | 176.11 | 177.64 |
| C3n.2n top | 4.493 | 197.15 | 197.26 |
| C3n.2n bottom | 4.631 | 211.15 | 211.36 |
| C3n.3n top | 4.799 | 231.89 | 233.15 |
| C3n.3n bottom | 4.896 | 242.89 | 244.14 |
| C3n.4n top | 4.997 | 255.63 | 255.89 |
| C3n.4n bottom | 5.235 | 279.39 | 280.14 |

fabric. In flow-aligned fabric the $I_1$ axes are oriented parallel to the mean flow direction, while in a flow-transverse fabric the $I_1$ axes are oriented perpendicular to the flow direction. In turbulent flows, grains settling from suspension tend to orient with their $I_1$ axes parallel to the flow direction and imbricated upstream (Rusnak, 1957; Allen, 1984). This flow-aligned orientation can be changed into a more stable flow-transverse orientation when the flow becomes strong enough to lift grains and roll them over the surface (e.g. Schwarzacher, 1963; Johansson, 1964; Hendry, 1976; Harms et al., 1982). In both cases the *foliation* planes can be imbricated dipping upstream (Harms et al., 1982) and the comparison of their orientation with $I_1$ axes can be used to recognize the flow-aligned and flow-transverse fabrics. Deviations from the flow-aligned or the flow-transverse fabrics can occur for a number of reasons which include spatial changes in current direction, bed surface irregularities, incomplete reorientation of a rolling fabric into a flow-aligned fabric or vice versa, changes in bed roughness and




*(caption on next page)*

**Fig. 7.** A and B) Equal area projection of the main anisotropy axis $I_1$ and foliation planes for the specimens sets with $P' \geq 1.1$ and $P' < 1.1$, respectively. $I_1$ axis are shown in different colours depending on their flow pattern. The set with $P' \geq 1.1$, mostly taken above 168 ± 2 m CSF-A, shows foliation planes imbricated along the current direction, in this case imbrications approximately toward NE and SW indicates currents flowing alternatively in these opposite directions. C and D) Current directions shows in the circular plots (dots) together with their rose diagram. The set with $P' \geq 1.1$ shows two distinct modal values while the set with $P' < 1.1$ have uniformly distributed directions. E) Von Mises distributions and mean values (red arrows) for the set of current directions with $P' \geq 1.1$. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

post-depositional modification by bioturbation or soft-sediment deformation (e.g., Baas et al., 2007 and references therein).

We recognize the pattern of each specimen by comparing the angle (θ) between the direction of the magnetic lineation $I_1$ and that of the foliation plunge. If $\theta < 35^\circ$ the pattern is flow-aligned and the flow is taken equal to declination of the $I_1$ axis in the direction of the foliation imbrication; if $\theta \geq 55^\circ$ the pattern is flow-transverse and the flow is the declination of $I_1$–90° in the direction of the foliation imbrication. The intermediate case ($35^\circ < \theta \leq 55^\circ$) is handled by taking directly the imbrication direction of the foliation plane as the flow direction.

In Site U1467, we found that the AIRM is large enough to produce a well-defined pattern of orientations only if the degree of anisotropy $P' \geq 1.1$, which mostly comprises specimens with flow-transverse pattern and located in the upper part on the sediment column. Current directions, foliation planes and $I_1$ directions are shown in Fig. 7 in separated sets for $P' \geq 1.1$ and $P' < 1.1$. In the set with $P' \geq 1.1$, the current directions fall into two distinct groups with nearly opposite modal directions highlighted by the rose diagram (Fig. 7c). Foliation planes also have the opposite plunge and their direction is consistent with the current modes (Fig. 7a). The mean current directions are computed as a mixture of 2 Von Mises distributions, which is necessary since we have two groups of directions and Von Mises distributions are unimodal. Calculations were performed using the R-package "movMF" (Hornik and Grün, 2014) and returned two independent distributions, the first with mean direction m = 45.3° and concentration parameter k = 6.3, and the second with mean direction m 227.4° and concentration parameter k = 2.2 (Fig. 7e). The current directions are nearly antipodal as expected for seasonally reversing monsoon-driven currents. In the set with $P' < 1.1$, the flow directions, the foliation planes and $I_1$ axis appear dispersed, probably because bottom currents were absent or too weak to produce a coherent directional pattern in elongated sediment particles (Fig. 7b and d). A Kuiper test for uniformity accepted the Null hypothesis at the 95% confidence level testifying that these directions do not have a preferential orientation. According to these observations stratigraphic intervals with larger P′ indicate the presence of stronger bottom currents that flow alternatively toward NE and SW. The N-S components of the observed currents are interpreted as a deflection of equatorial zonal currents in the Inner Sea of the Maldives where bottom currents are forced to follow the sea floor morphology and the directions of the main channels. Inferred current directions are virtually identical to those of present-day bottom current data measured by acoustic Doppler profiler by Lüdmann et al. (2013).

The presence of bottom currents is not constant throughout the stratigraphic record. In fact the degree of anisotropy P′ is generally very small in the lower part of the stratigraphic column (mean 1.05 ± 0.03) and shows larger values (mean 1.22 ± 0.17) in the upper part with a sudden increase at about 168 ± 2 m CSF-A, which corresponds to the top of Chron C3n.1n and an age of about 4.2 Ma (Fig. 8). The increase of anisotropy in the upper 168 m CSF-A is synchronous with a more gradual decrease of IRM intensity, which is indicative of a decreased concentration of magnetic minerals. The decrease of IRM intensity can be interpreted as a superimposed long-term trend with an acceleration starting at the depth of ~168 m CSF-A (Fig. 8b). No changes in the main lithological units were observed at this depth (Betzler et al., 2017), however the decreased concentration of magnetite is followed by deteriorated quality of the paleomagnetic measurements and decreased sedimentation rate in the upper part of Site U1467. From the sedimentological point of view, the decrease of IRM is interpreted as a consequence of changes in the sediment transport mechanism -controlled by wind driven currents- that transferred the sediments and the single-domain magnetite, possibly of biogenic origin, from the shallow platform to the deeper water of Site U1467 (Lüdmann et al., 2013). This process is modified by the increased monsoon strength starting at ~168 m CSF-A and the depocenter of drift deposits moving downstream. Regardless of the reason for the IRM decrease, the increased anisotropy can be associated with changes in sedimentation dynamics that lead to drift deposition and that has been related to the onset of strong modern monsoon system (Betzler et al., 2016b).

Our results suggest that starting from the lower Pliocene (ca. 4.2 Myr ago) the monsoon-related bottom currents became strong enough to significantly increase the degree of anisotropy and create a mostly transverse pattern in the sediments with large AIRM. Increased monsoon strength could qualitatively be explained with the onset of the intertropical convergence zones (ITCZ) to their present-day position. This implies a southern shift of the ITCZ south of the Himalayas and an increase in the latitudinal separation of the summer and winter ITCZ that moved the winter ITCZ south of the Maldives (e.g., Allen and Armstrong, 2012 and references therein). The Himalayas and Tibet have a primary influences on atmospheric circulation patterns and hence climate of the region. For this reason the surface uplift history of the Himalayan-Tibetan orogen has been suggested to be closely linked to the development of the Asian monsoon (Clift et al., 2008) and in fact, Tibetan plateau and Himalayan uplift is considered necessary for the presence of the strong present day monsoon (Prell and Kutzbach, 1997).

During the late Cenozoic the regional uplift may have occurred in two stages, one beginning in the Late Miocene, which probably led to the beginning of the drift deposition at 12.9 Ma (Betzler et al., 2016b), followed by a later Pliocene phase dated approximately from 5 to 2 Myr ago (Harrison et al., 1992; Zheng et al., 2000; An et al., 2001) that could have been recorded in Site U1467. Independent evidence supporting a coeval increase of monsoon intensity through enhanced precipitation, occurring at about 4 Ma, is given by the magnetic susceptibility record from ODP site 758, (Prell and Kutzbach, 1997; An et al., 2001), which is interpreted as the sea-level-mediated fluvial transport from the Ganges and other river systems draining the southern side of the Himalaya-Tibet plateau. Moreover, Zheng et al. (2000) interpret the increase in sedimentation rate and change in depositional facies from redbeds to upward-coarsening conglomerate and debris-flow deposits at the foot of the Kunlun Mountains as evidence for the uplift of the northwestern Tibetan Plateau between 3.5 and 4.5 Ma. The timing of increased current strength in the Maldives platform is compatible with the beginning of the Pliocene uplift stage, and in fact this could mark precisely the beginning of climatic influence of the Pliocene Himalayan uplift at 4.2 Ma.

## 5. Conclusions

Paleomagnetic study of IODP Site U1467 provides a magnetic stratigraphy that gives an improved age model of the Pliocene portion of Site U1467 compensating for the scarcity of the biostratigraphic data in this time interval. This new age model can potentially be the basis for further astrochronological studies.

The analysis of the AIRM has shown evidence of bottom currents with alternating directions similar to the present-day currents. We found that the strength of the bottom currents inferred from the AIRM-corrected anisotropy degree P′ increased suddenly at about 4.2 Myr

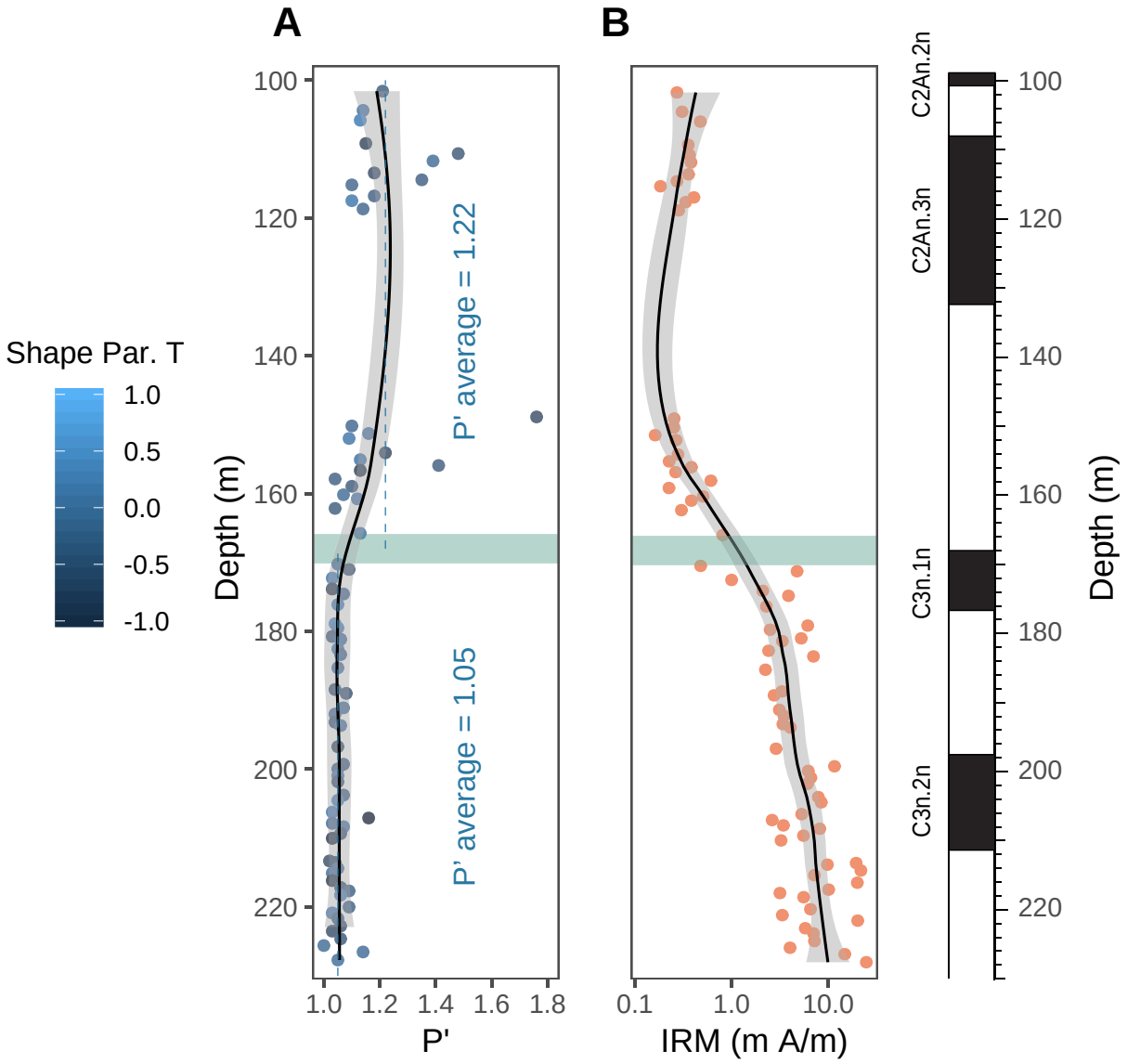



**Fig. 8.** Summary of anisotropy of isothermal remanent magnetization data versus depth. P′ indicates the corrected anisotropy degree, the shape parameter T is illustrated with a colour code. The IRM is indicative of concentration of magnetic minerals. Data have been smoothed using the locally weighted regression method (Cleveland, 1979; Cleveland et al., 1992) to illustrate the main trend. The 95% confidence level is shown by the grey band. The reversal polarity column provides a time frame and ties the age of the green band marking the shift toward higher anisotropy and lower IRM intensity to the top of chron C3n.1n. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

ago. This is interpreted as the formation of stronger equatorial undercurrents as a consequence of increased monsoon strength.

A number of studies relate the strength of Asian monsoon to the uplift of the Himalayas and Tibetan plateau. We observe that the timing of the increase of bottom currents (4.2 Ma) coincides with the increase of fluvial transport to the Bay of Bengal and is compatible with the beginning of the Late Pliocene phase of Himalayan uplift, suggesting that is represents the Maldives record of the Late Pliocene uplift phase. In this case our age model gives a precise timing of this event.

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